# On the Threshold Sea Surface Temperature for Onset of Convection

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- Notion of threshold SST based on strong deep convection zones for SST ≥ 27.5 C (Graham & Barnett,1987)
- Threshold is relative. Pattern ≈ same if add 2C everywhere (Randall et al 2001)
- Can we make concept more precise?
- Spatial dependence?
- Add local –/+ SST anomaly until convection stops/starts
- Experiments in QTCM: Quasi-equilibrium Tropical Circulation Model (> 500 ensemble experiments)

# Example of regions for which local threshold will be determined



## **Precipitation vs SST for various regions for DJF**

#### **Defining the threshold:**

#### Use 0.2 mm/day for onset of convection in ensemble avg.

(defining threshold between medium and large convection more arbitrary)

Threshold values: 297.2 to 299.4

#### **Realistic case:**

- SST = observed clim. + local anomaly
- Land temperature is modeled



### **Effect of box size**

- Western Pacific warm pool, all centered at (169E, 2S)
- Region where SST is lowered is 39°x19°, 28°x19°, 17°x1°1 and a 17°x11° in a 28°x19° area, respectively
- SST<sub>crit</sub> not very sensitive within this range



## Precip. & Winds (850mb) at SST<sub>crit</sub>

- Idealized SST climatology case
- SST region bounded by 157.5E - 147.375 X 5.625S - 5.625N
- SST<sub>crit</sub> = temp at which prec = 0.2mm/day





# **T** Climatology and Anomaly

- Idealized SST climatology case
- Anomaly box area bounded by 157.5E - 147.375E × 5.625S - 5.625N
- T denotes tropospheric vertical average
- T anomaly amplitude relatively small



## Idealized SST climatology case - SST, $SST_{crit}$ , $\hat{T}$ , $\hat{T}_{c}$



# Box area avg. SST vs. **T**, **T**<sub>c</sub> Prec, **q**

- Idealized SST clim
- Specific humidity q tends to drop with local SST
- $\hat{T}_c$  drops with q
- **T** drops more slowly
- CAPE ~  $\hat{T}_c \hat{T}$
- Precip drops with CAPE
- But variations within box so point with max CAPE controls onset



# Idealized SST +/-2C - SST, SST<sub>crit</sub>, $\hat{T}$ , $\hat{T}_c$

- SST<sub>crit</sub> is relative to surrounding regions
- Adding ±2C to all SST just shifts the entire pattern (approximately)



# NoLand-DJF & DJF (realistic case) SST, $SST_{crit}$ , $\hat{T}$ , $\hat{T}_{c}$

#### NoLand case:

- SST<sub>crit</sub> differs by > 1C between central Pacific and Atlantic or Indian (2C to E. Pac.)
- **T** smoother and shifted eastward rel. to clim SST
- $SST_{crit}$  tends to follow  $\hat{T}$
- Approx. consistent with **T**
- Postulate T̂<sub>c</sub> drawn toward local SST, T̂ strongly affected by large-scale wave dynamics

#### **Realistic case:**

• Similar but with exceptions ? (e.g., Atlantic)



#### Realistic case - DJF, MAM, JJA, SON - SST, SST<sub>crit</sub>, T, & T<sub>c</sub>

- SST<sub>crit</sub> changes with season as well as region
- Hard to find simple rule for SST<sub>crit</sub> relation to clim. SST
- Regional and land effects changing with seasonal climatology



### **Summary**

- **Defining SST**<sub>crit</sub> by local SST anomaly experiments and low onset threshold works in deep tropics
- SST<sub>crit</sub> is relative (to other regions): a very large scale SST change has completely different effects (e.g., little effect on precipitation pattern) than regional SST anomalies
- SST<sub>crit</sub> has substantial spatial and seasonal variations (smaller than range of SST climatology)
- For idealized SST climatology or No-Land case (with observed/interpolated SST) the relation of SST<sub>crit</sub> to SST can be understood by two effects on CAPE: wave dynamics on  $\hat{T}$  and local SST on q
- For realistic case the local effects involve advection, relation to neighboring land regions etc.  $\Rightarrow$  SST<sub>crit</sub> bears complex relation to SST climatology





