Dynamical versus Land-Surface Factors in the African Monsoon

J. David Neelin, UCLA

Collaborators: Chia Chou, Academia Sinica, Taiwan; Hui Su, UCLA; Ning Zeng, U. Maryland

•Dynamical and land-surface factors in summer monsoon climatology

Contrast of N. Amer. and Asian cases with African case
 * Experiments in an intermediate complexity model

•The question of land-surface feedbacks in teleconnected SST impacts in interannual-interdecadal variability

Climate Systems Interactions Group

http://www.atmos.ucla.edu/~csi

Factors in summer monsoon extent

>Thermodynamic:

•positive net flux (Rad +
SH + Latent) into
atmospheric column =
TOA over land

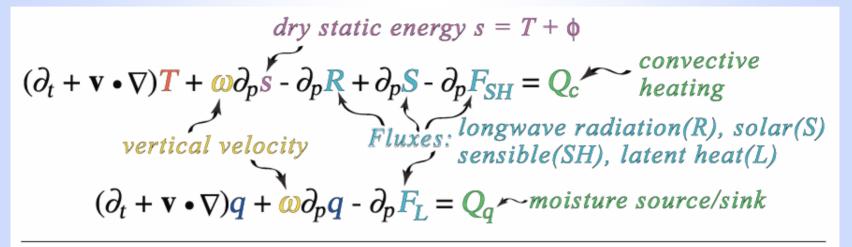
≻Land-Surface processes:

- •Albedo
- •Soil wetness/ evapotranspiration
- Ocean heat storage and transport

≻Dynamic:

- •"ventilation mechanism"
 - $\mathbf{v} \cdot \nabla(q + T)$ importing low moist static energy air
- wave dynamics
 - Kelvinoid
 - Rossby-related "interactive Rodwell-Hoskins" mechanism

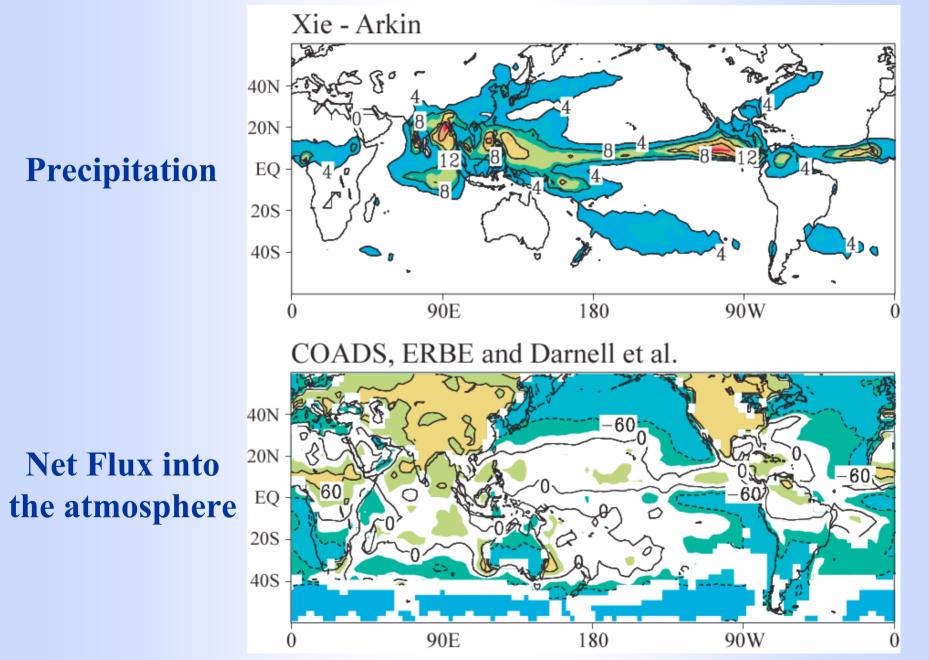
Temperature *T* **and Moisture** *q* **equations**



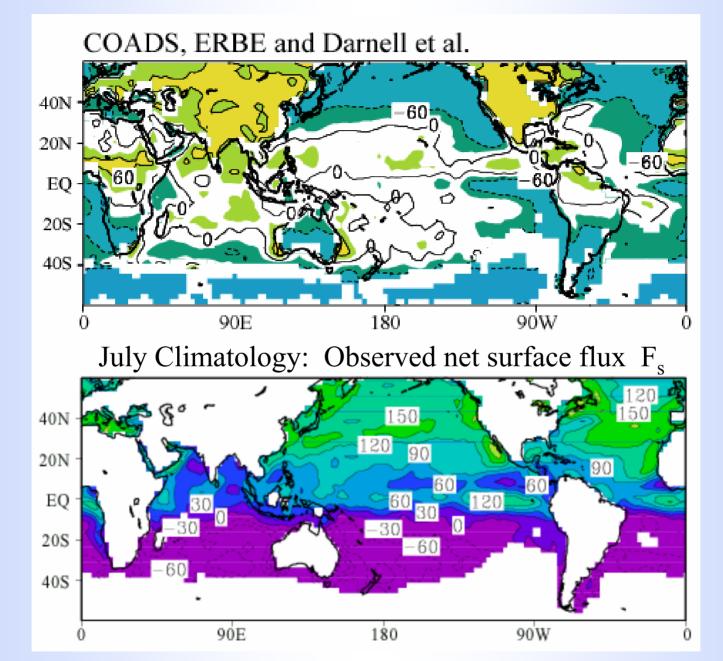
Energy constraint in vertical integral $\langle \rangle$ $\langle Q_c \rangle = -\langle Q_q \rangle$

 $\frac{\langle \text{Moist static energy equation} \rangle}{\langle (\partial_t + \mathbf{v} \cdot \nabla)(\mathbf{T} + q) \rangle + \langle \omega \partial_p h \rangle} - F_{net} = 0$ $\frac{\langle \text{Transport of moist static}}{\langle \text{Net energy flux} \\ \text{energy by divergent flow}} \quad \text{Moist static energy flux} \\ \approx (\text{measure of divergence}) \\ \times \text{ gross moist stability}} \quad h = s + q$

Observed climatology July



Observed net flux into atmosphere and net surface flux



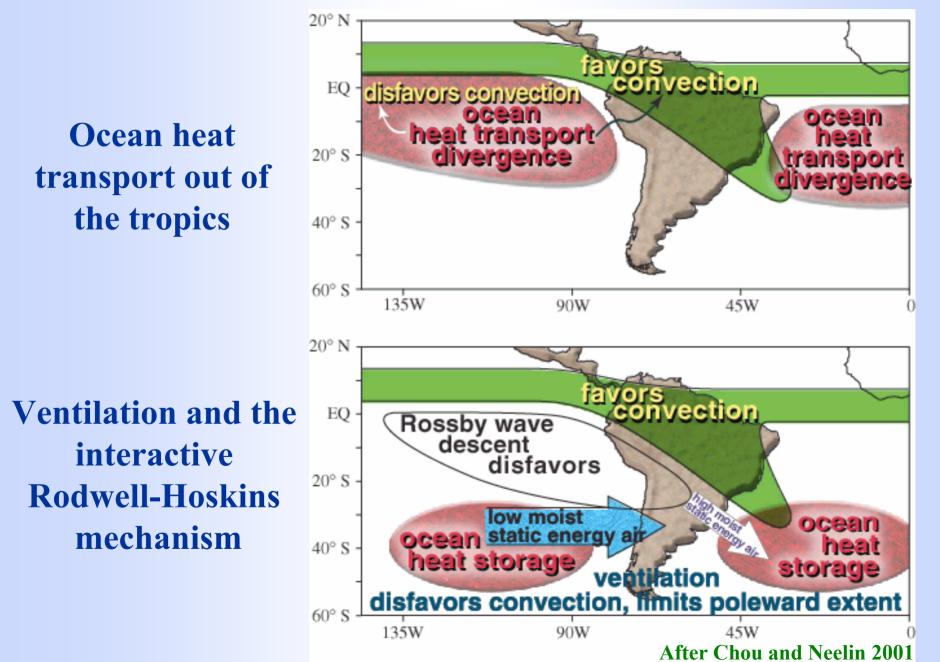
The "ventilation mechanism"

- import of low moist static energy air from ocean where heat storage opposes summer warming
- oceanic air: cooler and moisture is lower than convection threshold over warm continent
- **import** to continents by wind (including upper level jets) via advection terms in temperature and moisture equations

The "interactive Rodwell-Hoskins mechanism"

- Rodwell and Hoskins (1996): imposed convective heating in Asia gives Rossby wave descent pattern to west, enhancing deserts.
- when convection is interactive: associated flow feeds back on heating, creating characteristic convection/dry region pattern
 - » we emphasize feedback (convection ⇔ baroclinic Rossby wave dynamics), hence:
 - » "interactive Rodwell-Hoskins" (IRH) mechanism

Mechanisms affecting convective zones (S. American case)



Quasi-equilibrium Tropical circulation model:

- Primitive equations projected onto vertical basis functions from convective quasi-equilibrium analytical solutions
- for Betts-Miller (1986) convective scheme, accurate vertical structure in deep convective regions for low vertical resolution
- GCM-like parameters but easier to analyze

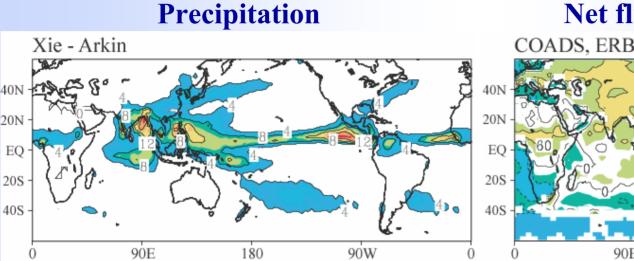
Radiation/cloud parameterization:

- Longwave and shortwave schemes simplified from GCM schemes (Harshvardhan et al. 1987, Fu and Liou 1993)
- deep convective cloud, CsCc fraction param. on precip

Simple land model:

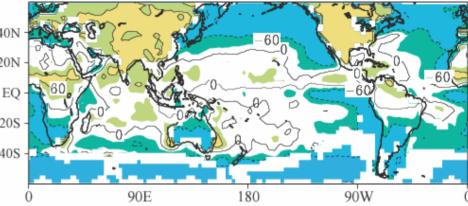
- 1 soil moisture layer; evapotranspiration with stomatal/root resistance dep. on surface type (e.g., forest, desert, grassland)
- low heat capacity; Darnell et al 1992 albedo
- http://www.atmos.ucla.edu/~csi/QTCM

Observed climatology July



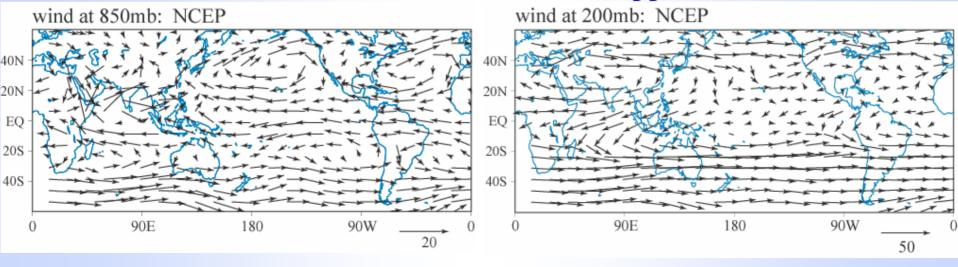
Net flux into atmosphere

COADS, ERBE and Darnell et al.



Low-level wind

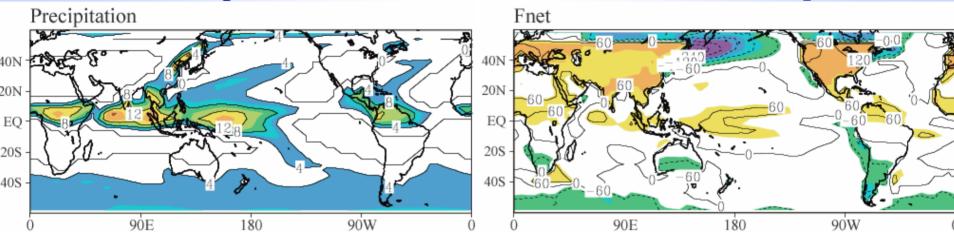
Upper-level wind



QTCM climatology July (coupled to a mixed-layer ocean)

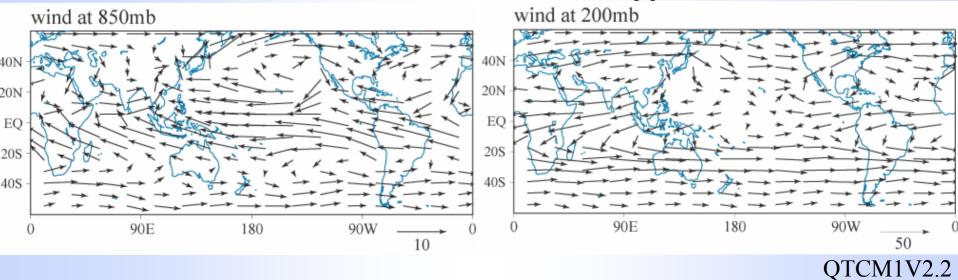
Precipitation

Net flux into atmosphere



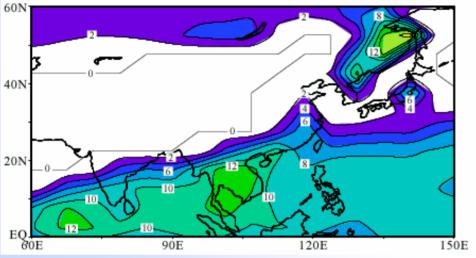
Low-level wind

Upper-level wind

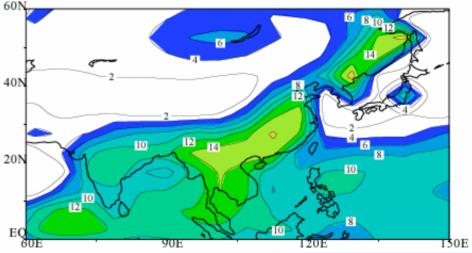


Asian region case – July Precipitation

Control

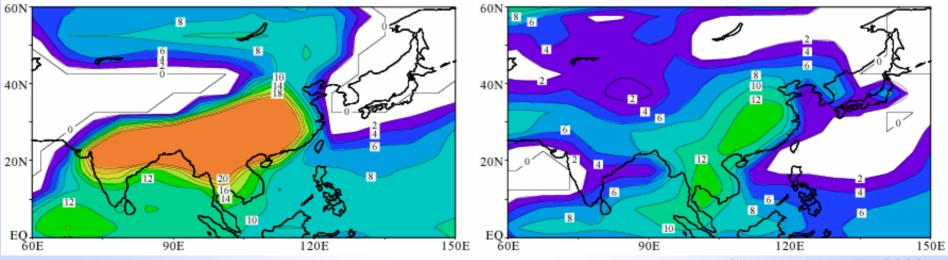


Saturated soil moisture



No ventilation: $v \bullet \nabla q$, $v \bullet \nabla T$ set to zero

No β -effect: f = constant

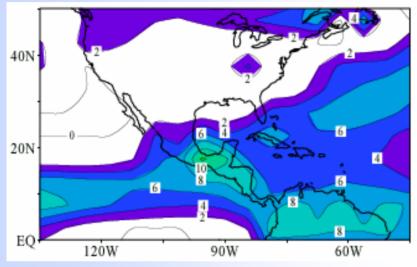


Chou and Neelin 2003

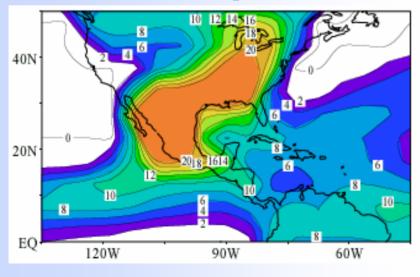
North American region case

July Precipitation

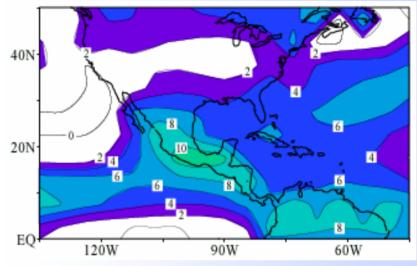
Control



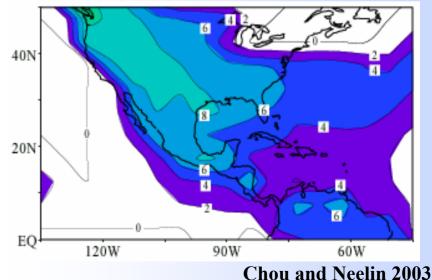
No ventilation: $v \bullet \nabla q$, $v \bullet \nabla T$ set to zero



Saturated soil moisture



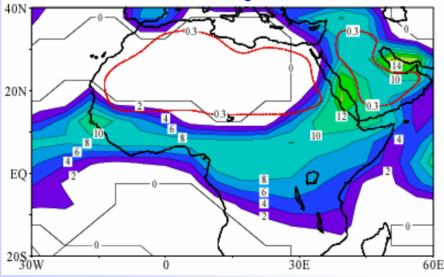
No β -effect: f = constant in region



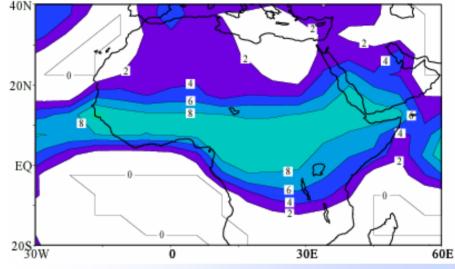
African region case (observed albedo) July **Precipitation** Control

40N 20N ź 6 4 26 EQ 208 AW 30E 60E

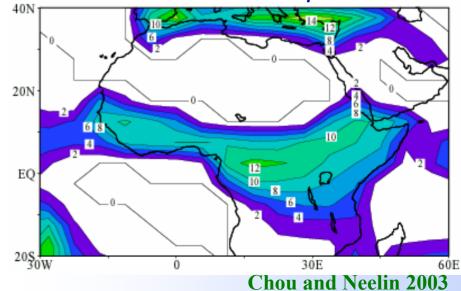
No ventilation: $v \bullet \nabla q$, $v \bullet \nabla T$ set to zero



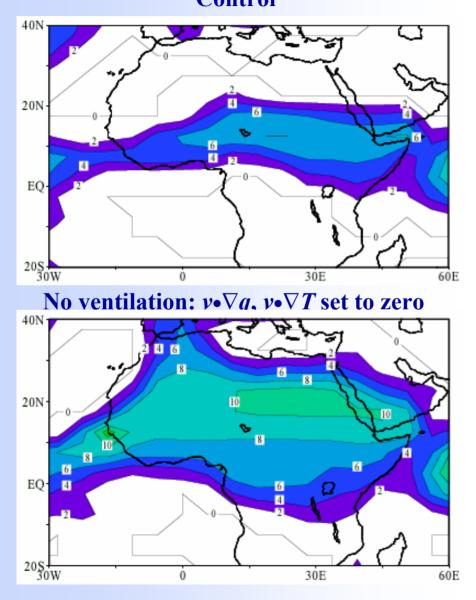
Saturated soil moisture



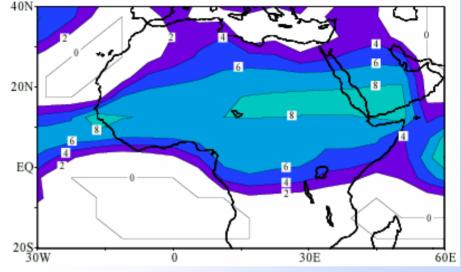
No ventilation and no β-effect:



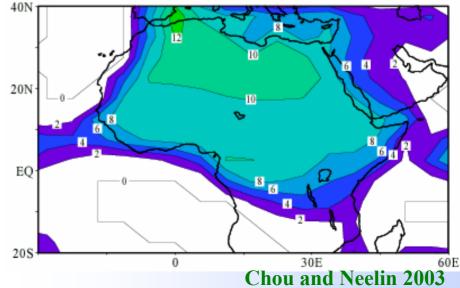
African region constant albedo case (0.26 over Africa) July Precipitation Control



Saturated soil moisture

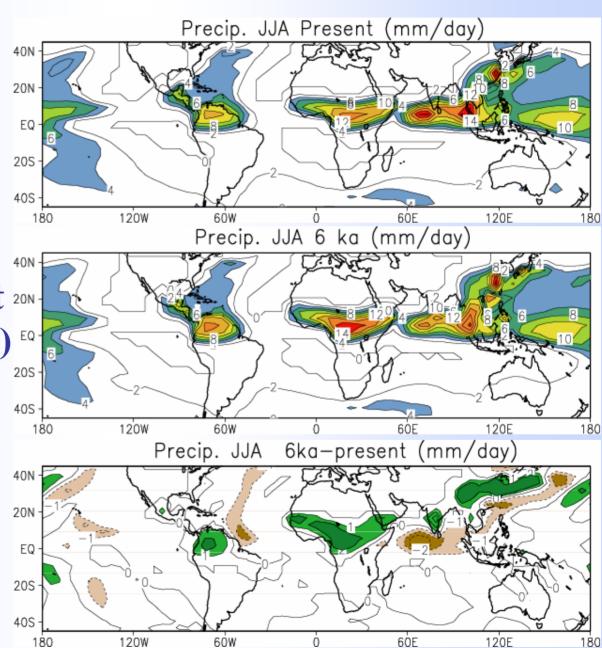


No ventilation and no B-effect:



QTCM PMIP - type 6ka BP simulation

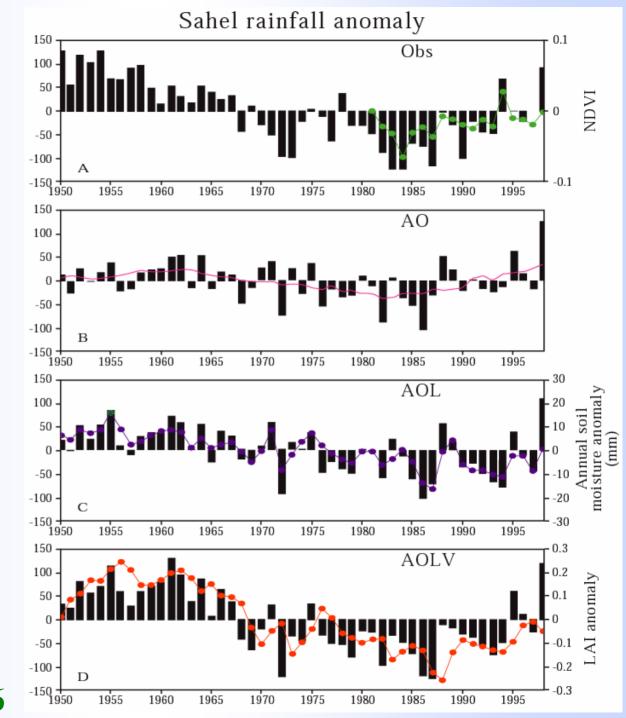
- For comparison to Paleo Model 205-Intercomparison 405-18
 Project (PMIP) mid-Holocene experiment 201-Joussaume et al (1999) EQ-
 - Present day SST, albedo
 6kaBP orbital params., CO2



Drying trend in the Sahel 1950s -1980s

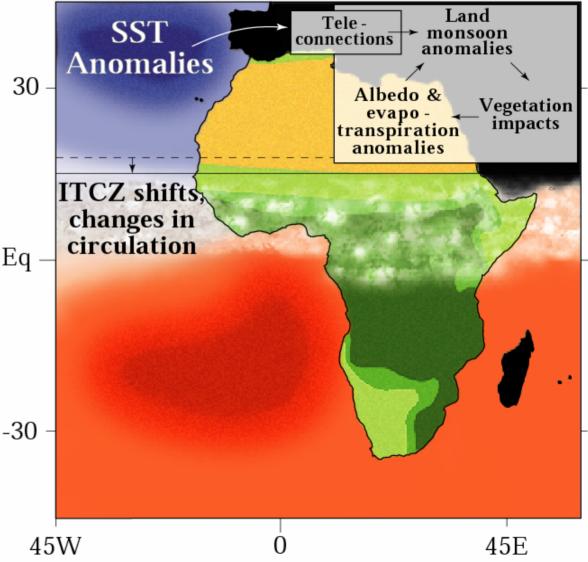
(A) Observed
(B) Forced by SST
(C) Amplified by
soil moisture and
(D) vegetation
Zeng et al 1999

- Land-Surface
 e.g., Charney 1975
 Xue and Shukla 1993
 Nicholson 2000
- SST impact e.g., Folland et al 1986



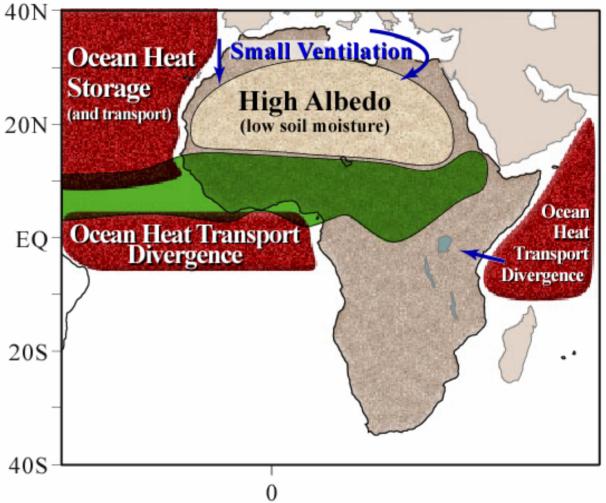
Interdecadal - interannual variability

- Teleconnected SST impacts, possibly enhanced by landsurface feedbacks
- Dynamical factors
 apparently dominate Eq
 (exact mechanisms
 TBD)
- Variability tends to smooth gradient in precipitation climatology



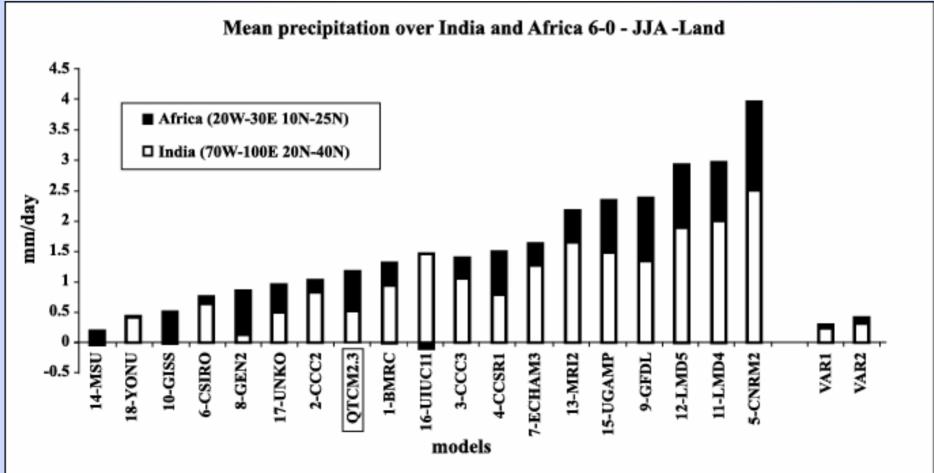
African northern summer monsoon climatology

- Albedo is leading effect on poleward extent of convection ²⁰
- Dynamical mechanisms:
 - * affect margin of convective zone
 * take over if alked
 - take over if albedo gradient is flattened



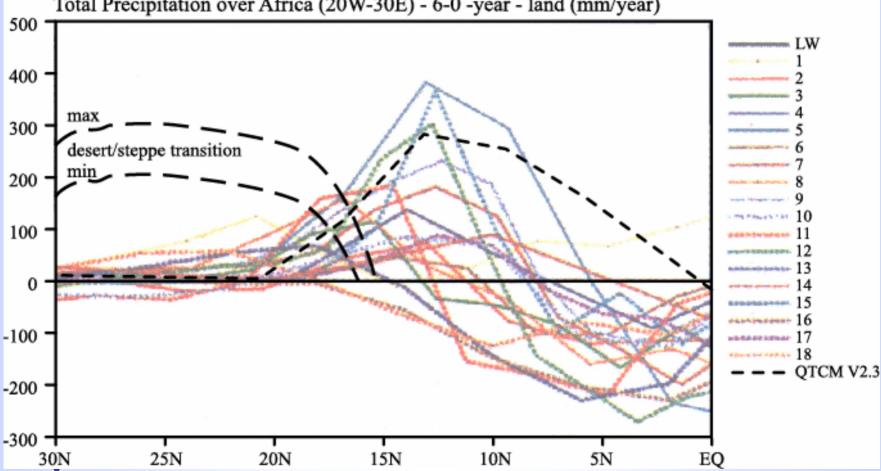
PMIP 6kaBP precipitation change over India & Africa

After Joussaume et al 1999



PMIP 6kaBP precipitation change over Africa

After Joussaume et al 1999



Total Precipitation over Africa (20W-30E) - 6-0 -year - land (mm/year)